

# Using structure from motion and high-resolution digital elevation models to investigate the relationships between emplacement history and lava surface roughness: Hawaii, California and Mars

D. H. James (davita.james@unco.edu) and S. W. Anderson (steven.anderson@unco.edu)  
University of Northern Colorado (Campus Box 100, 501 20th St, Greeley, CO 80639)



## Introduction

We can infer eruption conditions using variations in roughness at different scales on a lava flow. Mauna Ulu, Hawaii and Amboy, California were used as analogues for a range of martian lava flow surfaces.

- Hawaii offers an opportunity to observe young flows, but the weathering and erosional processes are significantly different from those on Mars<sup>1</sup>.
- Lava flows at Amboy are older than those produced by Mauna Ulu, and display varying levels of mantling by wind-blown sand similar to expectations for Mars.

This research aims to answer several questions about lava surface roughness and geologic history. The one presented here is:

- How do roughness values of mantled lava flows on Mars compare to values at Amboy?

## Methods

- We selected sites located on and off the wind streak at Amboy (1m/pixel, airborne LiDAR) to compare the effect of mantling on lava surface roughness there with surfaces on Mars (2m/pixel, HiRISE)
- Roughness Doughnut (RD) method displays roughness values relative to the focus point, based on neighborhood elevation statistics.
- RD considers cells along the circumference of a circle, unlike the Topographic Position Index which averages the full area of the circle, thereby dampening roughness signals. RD values and related elevation products should highlight partially mantled surfaces
- Neighborhood size was adjusted to change the scale of roughness being observed.
  - The outer radius or 'doughnut thickness' remained constant at 2 cells.
- RD rasters were calculated by:
 
$$\frac{\text{mean raster} - \text{original DEM}}{\text{range raster}} = \text{RD}$$
- The larger the RD value, the lower the topographic variation at the selected scale. A small value represents the inverse of this – a wider variety of elevations at the scale.
- A Principal Component Analysis (PCA) was used to identify patterns in the data.

Figure 2: Location of sites relative to the crater and wind streak at Amboy, California



Figure 3: mantled lavas in the Tharsis region of Mars. Location of selected sites marked by yellow pin (HiRISE)

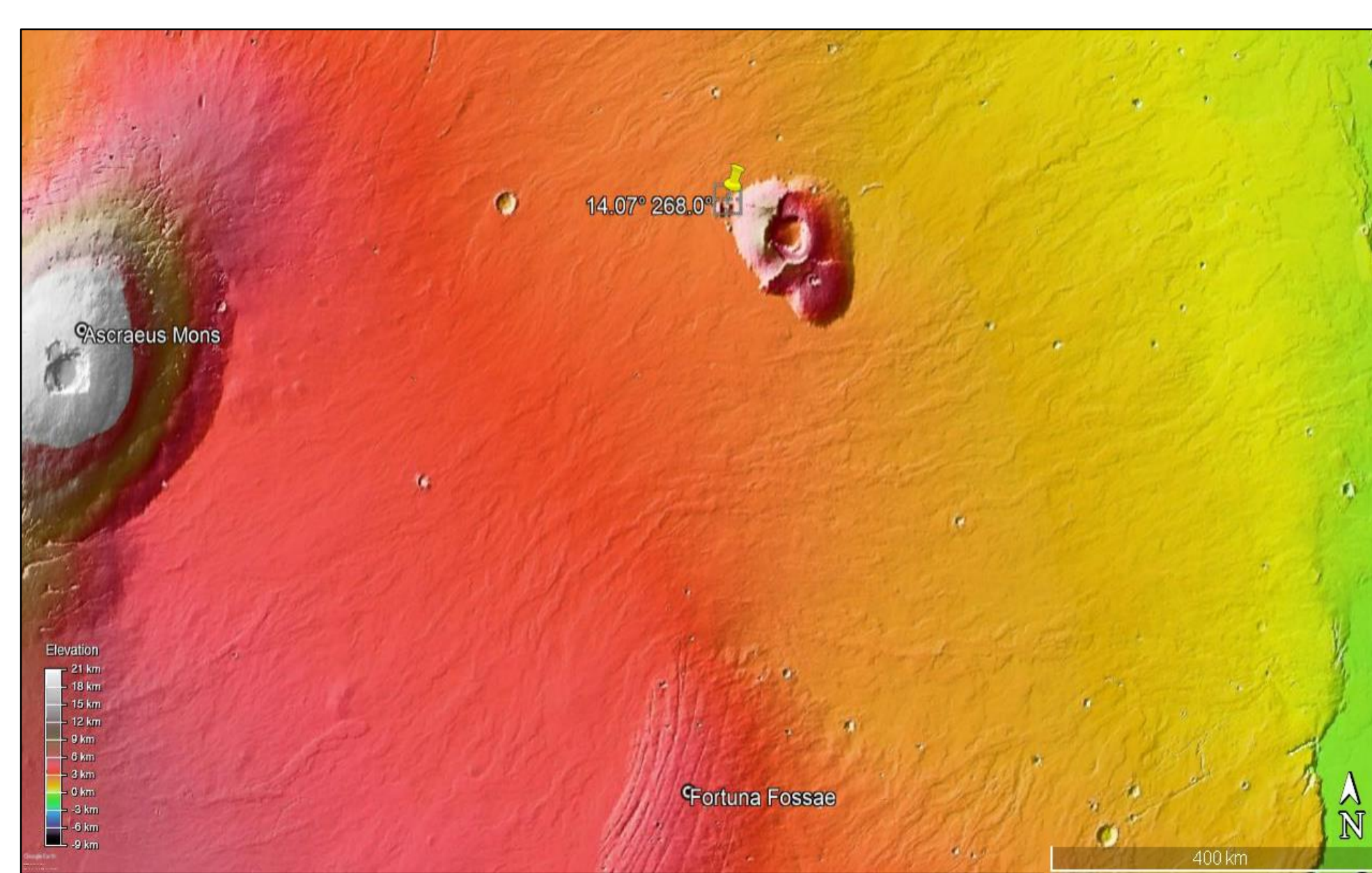
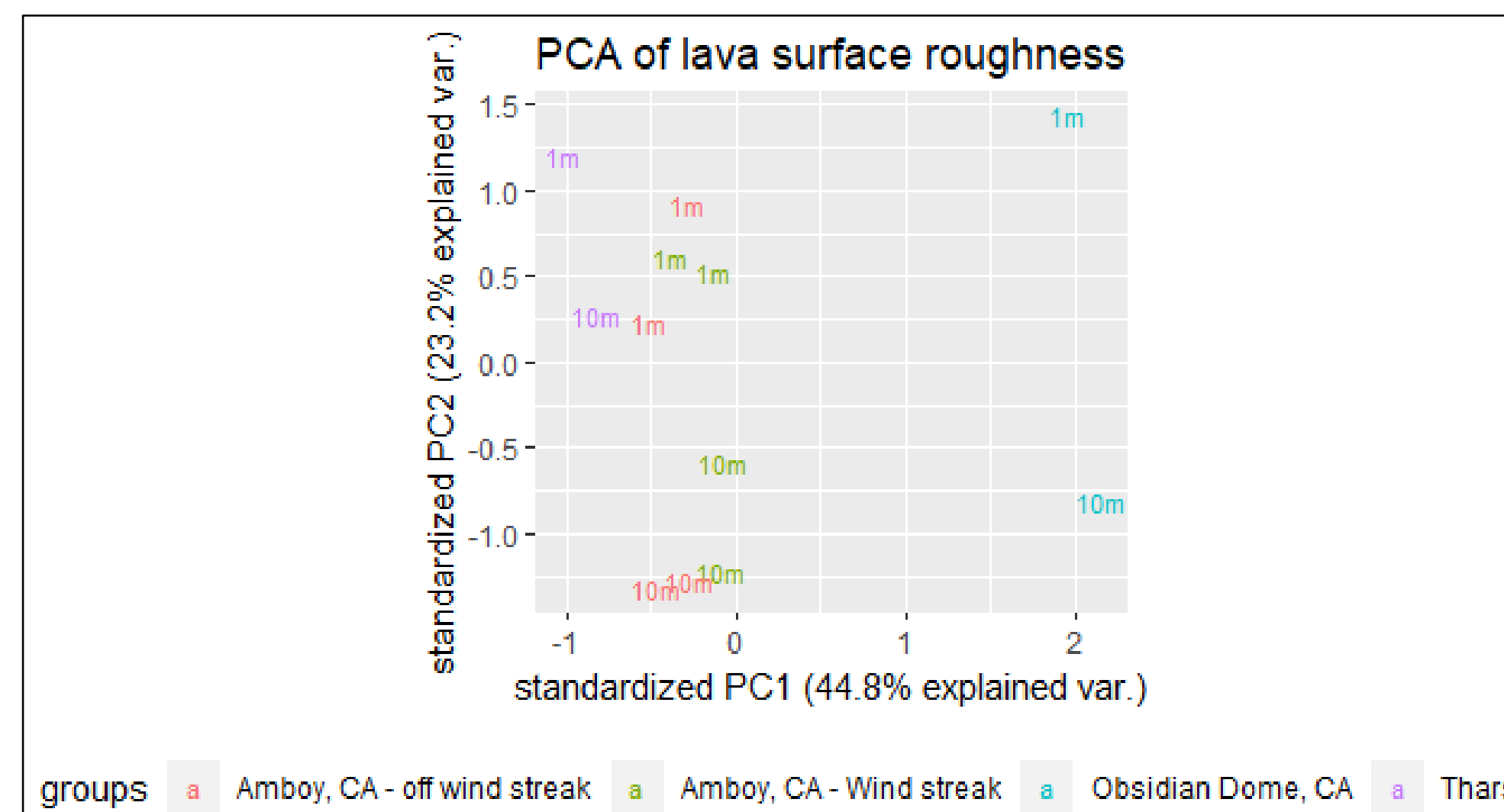


Figure 4: Principal Component Analysis of roughness data from Amboy, and Tharsis. Obsidian Dome was included to show that this method can be used to distinguish between mafic and silicic flows



## Results

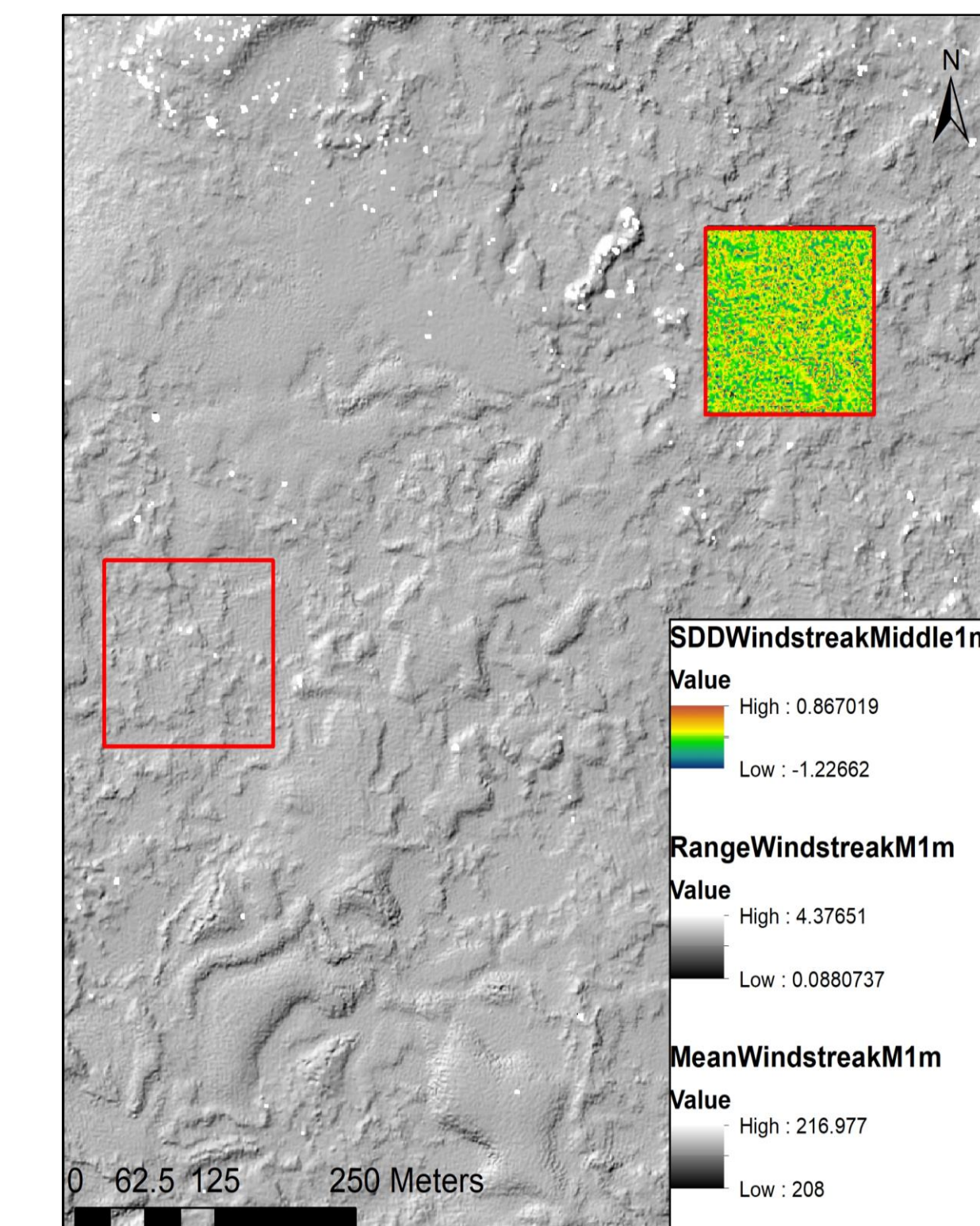
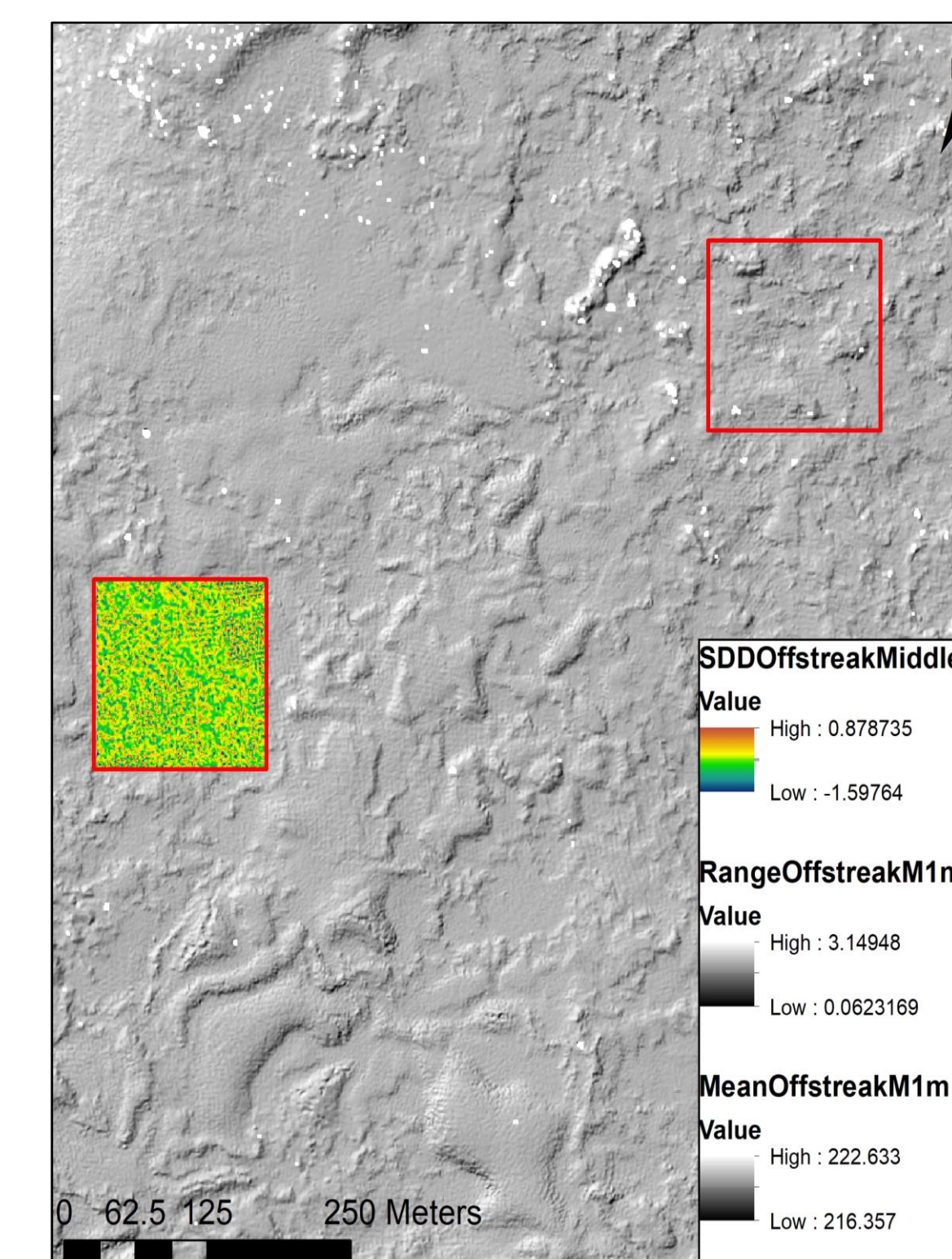


Figure 5a (left): Surface roughness at the 1 meter scale on an area outside of the wind streak at Amboy.  
Figure 5b (right): Surface roughness at the 1 meter scale on the wind streak at Amboy

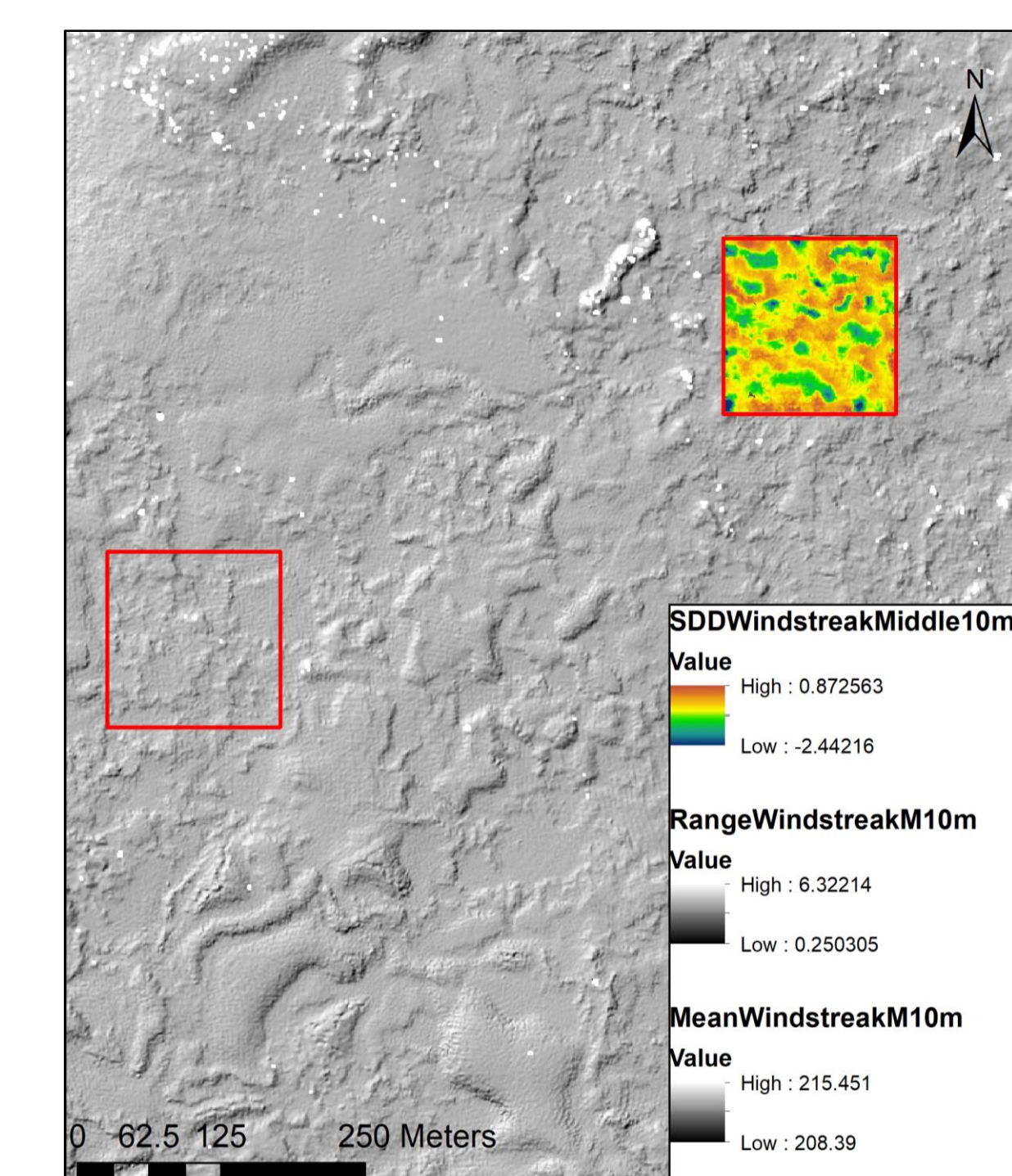
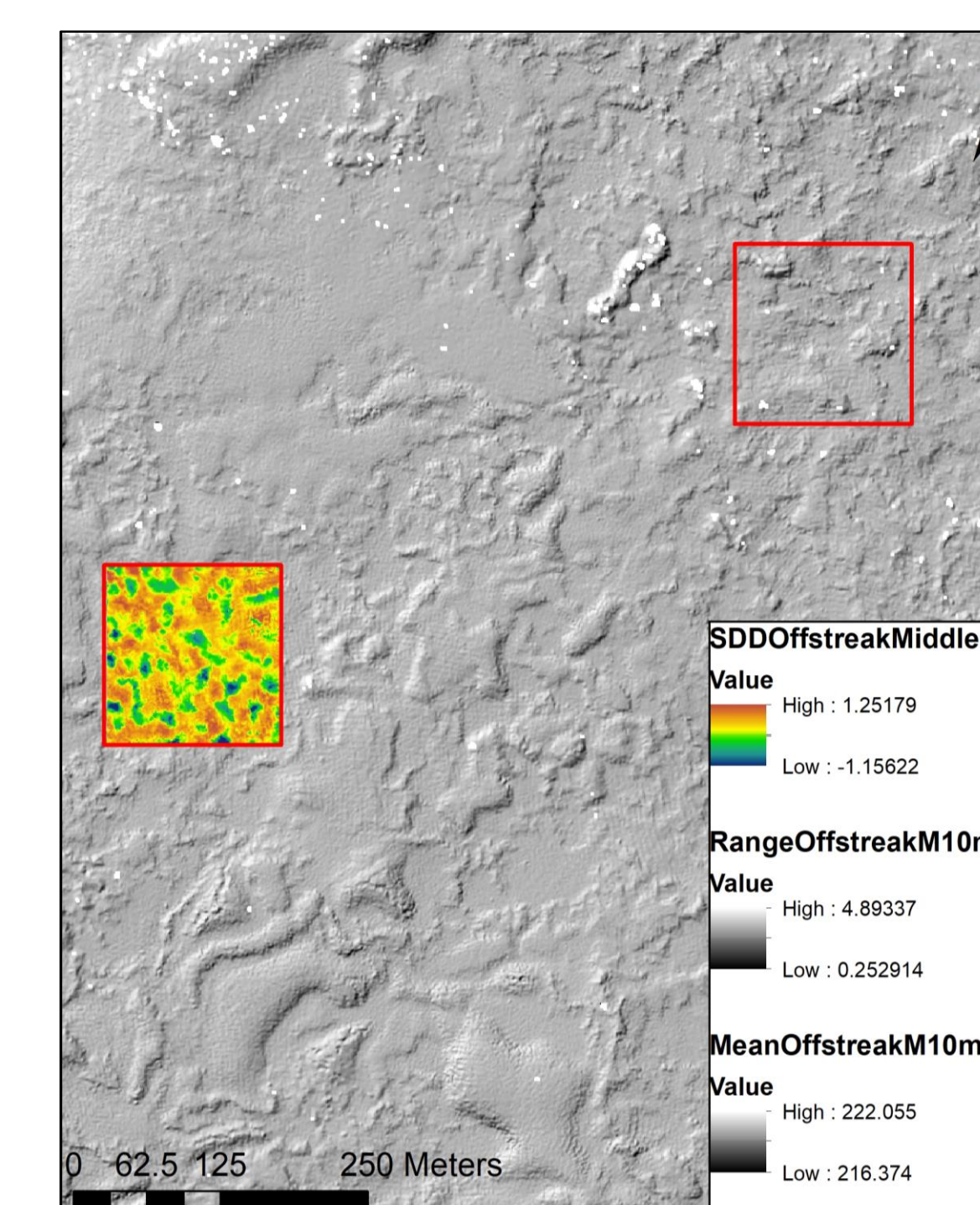


Figure 6a (left): Surface roughness at the 10 meter scale on an area outside of the wind streak at Amboy.  
Figure 6b (right): Surface roughness at the 10 meter scale on the wind streak at Amboy

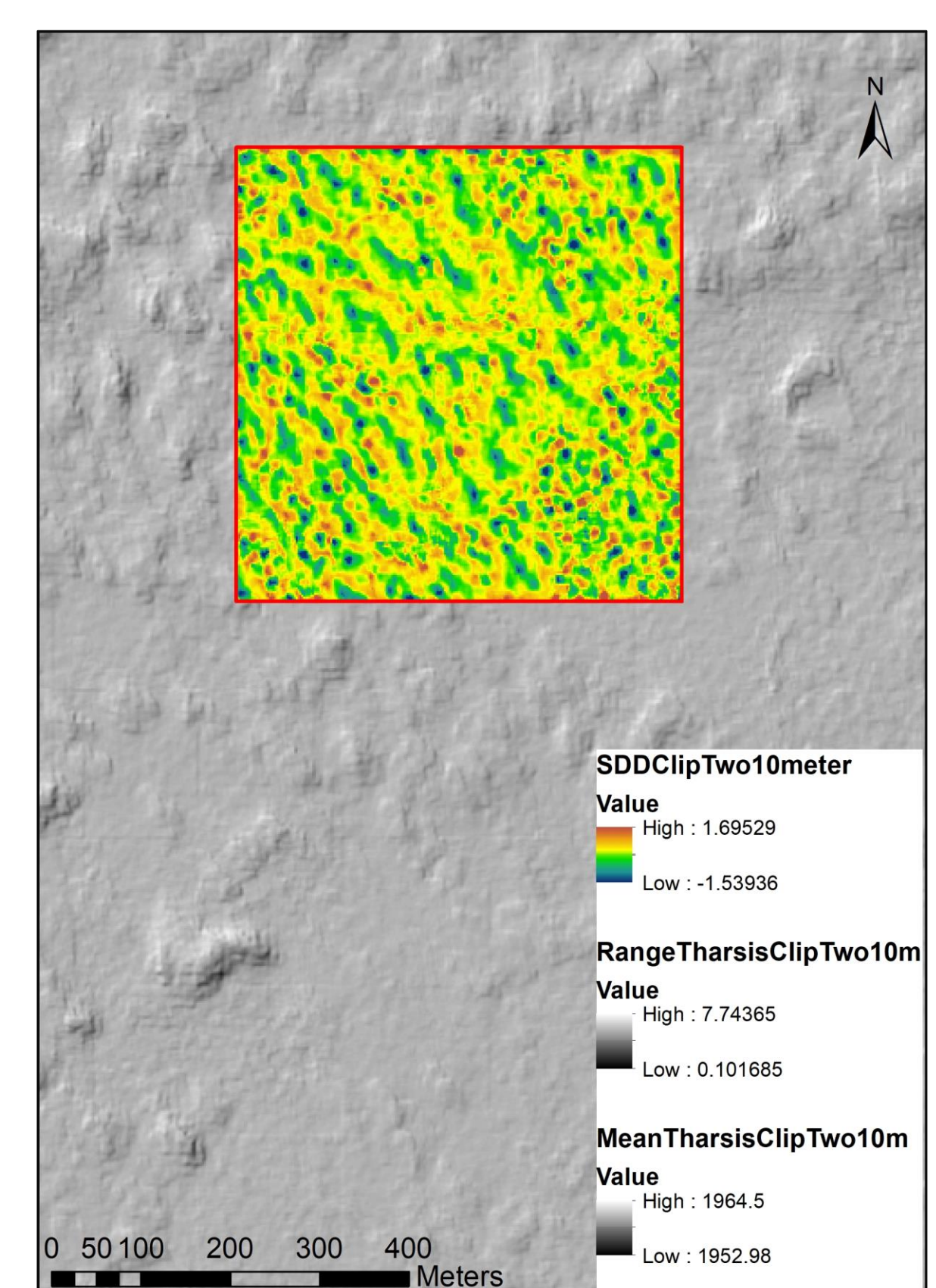
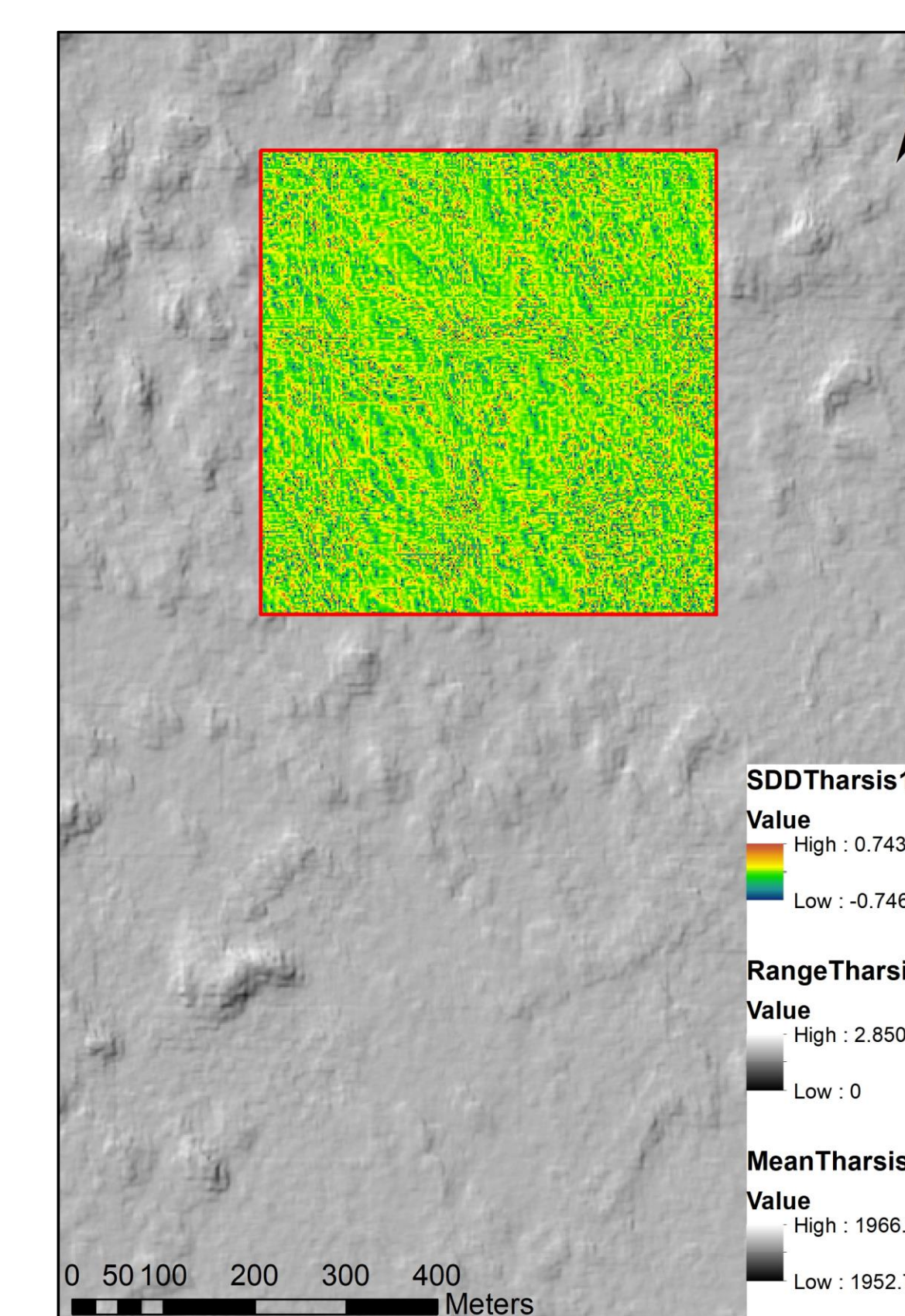
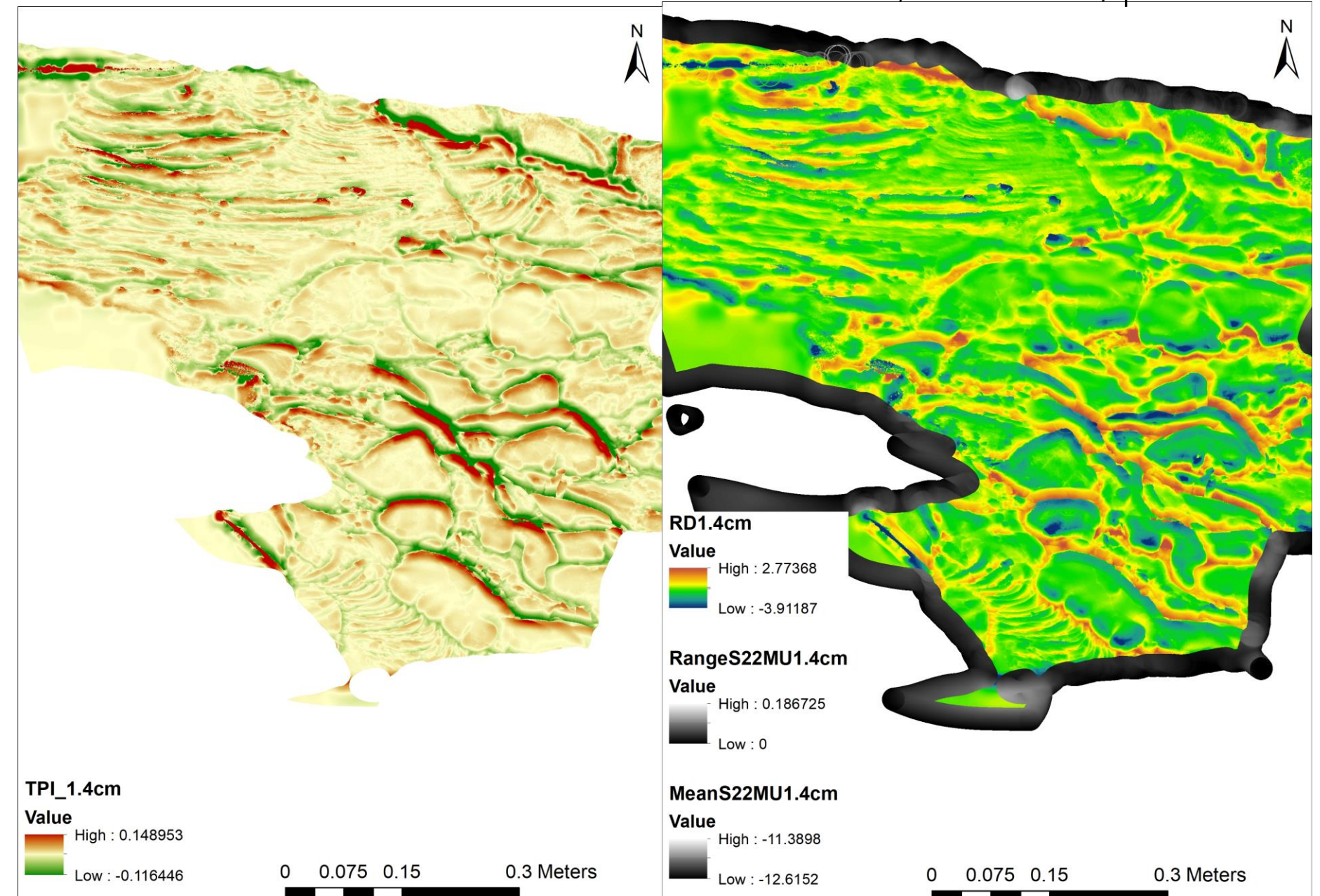


Figure 7a(left): Surface roughness at the 1 meter scale on an area in the Tharsis region.  
Figure 7b (right): Surface roughness at the 10 meter scale on an area in the Tharsis region

Figure 1: TPI (left) and RD (right) roughness comparison using a Mauna Ulu Structure from Motion DEM, 0.14 mm/pixel



## Conclusions

- RD values increase with the rate of mantling, though this trend is not robust
- As sand settles preferentially in the lowest regions, the elevation range as well as the final SDD value increase.
  - This is true for both comparison sites at Amboy, at both the 1-meter and 10-meter scales, as well as Tharsis
- Trends and patterns of roughness values are similar between locations, though the features on Mars are significantly larger than those in California.

## Acknowledgments

This work was funded by research and travel grants from the College of Natural and Health Sciences, Graduate Student Association and Earth and Atmospheric Sciences Department of UNC. Travel to the 50<sup>th</sup> Lunar and Planetary Science Conference to present this poster was supported by the Lunar and Planetary Institute.

## References

- Byrnes J. M. et al. (2004) JGR, 135, 169-193
- Mallonee H.C. et al. (2017) LPSC XLVIII abstract #2975.
- Finnegan D. C. et al. (2004) LPSC XXXV abstract #1736.
- Fink J. H. and Anderson S. W. (2000) in Encyclopedia of Volcanoes, 1, 307-320.
- Mallonee H. C. et al. (2017) LPS XLVIII, abstract #2992.
- Fan K.A. et al. (2018) LPSC LXIX abstract #2526.
- Tolometti G. D. et al. (2017) LPSC XLVIII abstract #1643.
- Zanetti M. et al. (2018) LPSC LXIX abstract #2361.
- Rogers A. D. and Head J. W. (2017) LPSC XLVIII abstract #1347.
- Tanaka K. L. (2000), Icarus, 144, 254-266.
- Guest, J. E., et al. (1987) Bull. Volc., 49, 527 – 540.
- Byrnes J.M. and Crown D. A. (2002) JGR, 107, 5079.
- Moncrief S. R. and Rowland S. K., (1991) LPSC XXII, 913-914.
- Philips F. M. (2003) Geomorphology, 53, 199-208.
- Chester-man et al. 1971.
- Byrnes et al. (2007) LPSC XXXVIII abstract #1908.
- Kienberger R. L. and Greeley R. (2011) 42nd LPSC, abstract #1053.
- Kreslavsky and Head (2000) JGR, 105, 695-711.
- Theilig E. and Greeley R. (1986) JGR, 91, 193-206.
- Kim J. R. and Muller J. P (2008) Intl. Archives of Photogrammetry, Rem. Sensing and Spatial Info. Sci. XXXVII 993-998.
- Whelley P. L. et al. (2014) Transactions on Geoscience and Remote Sensing, 52, 426-438.
- Simurda C.M. (2018) LPSC LXIX, abstract #2612.